TITLE:
Observations of Tidal Strain of the Earth by the Extensometer (Part II)

AUTHOR(S):
OZAWA, Izuo

CITATION:
OZAWA, Izuo. Observations of Tidal Strain of the Earth by the Extensometer (Part II). Bulletins - Disaster Prevention Research Institute, Kyoto University 1952, 3: 4-17

ISSUE DATE:
1952-12-15

URL:
http://hdl.handle.net/2433/123643

RIGHT:

京都大学学術情報リポジトリ
KURENAI
Kyoto University Research Information Repository

Observations of Tidal Strain of the Earth by the Extensometer (Part II)

By Izuo OZAWA

1. Horizontal linear strains of the earth caused by the earth-tide and by the crustal deformations are being observed in kyoto University by the use of Sassa-type invar-line extensometers with a sensibility of $10^{-8} \sim 10^{-9}$/mm. Some results of the observations at Osakayama[1] and Ikuno[1] have already been reported. In the present paper the tidal strains which were observed at the Makimine mine, Miyasaki Prefecture, and at Osakayama, were discussed more throughly than the first report, especially to the effects due to the oceanic tide.

2. Due to tidal generating potential W_2 of heavenly bodies the surface displacements S_r, S_θ, and S_φ along the radius, colatitude and longitude respectively of the earth are given as follows:

$$\left.\begin{aligned} S_r &= \frac{h}{g}\cdot W_2 \\ S_\theta &= \frac{l}{g}\cdot\frac{\partial W_2}{\partial\theta} \\ S_\varphi &= \frac{l}{g\sin\theta}\cdot\frac{\partial W_2}{\partial\varphi} \end{aligned}\right\} \tag{1}$$

where h is the Love-number, l the Lambret number, g acceleration of gravity, θ co-latitude and φ longitude. While the value of h is given to $h=0.60$[2] as the most probable value, deduced from Chandler's period of latitude variation and observations of tilting of the earth-tide. The Lambert's number l has been deduced heretofore from the relation

$$L=1+k-l \tag{2}$$

which is obtained from observation of the tidal change of latitude. According to Dr. Eiichi Nishimura,[3] $L=1.20$ and $l=0.07$ when k equals 0.267, as calculated from Chandler's[4] period. This value is for smaller than $l=\frac{3}{10}h=0.18$ which is a relation obtained when the earth is assumed to be an elastic sphere homogeneous and incompressible, although it was not possible to discuss this point because of insufficient of accurary of observations of latitude. On the other hand, Dr. Hitoshi Takeuchi[5] has obtained theoretically the value of $l=0.080 \sim 0.082$ using a model of the structure of the earth obtainable from the velocity of seismic waves inside the earth.

However from our observations of the change in horizontal linear strain of the earth due to the earth-tide with sufficient accuracy we obtained a most probable value of $l=0.05$ as reported in the previous paper. Taking the x-axis to the south and the y-axis to the west along the earth's surface, the components e_{xx}, e_{yy} and e_{xy} of horizontal linear strain will be expressed as follows:

$$\left.\begin{aligned} e_{xx} &= \frac{1}{a}\cdot\frac{\partial S_\theta}{\partial\theta} \\ e_{yy} &= \frac{1}{a\sin\theta}\cdot\frac{\partial S_\varphi}{\partial\varphi} \\ e_{xy} &= \frac{1}{a}\cdot\frac{\partial S_\varphi}{\partial\theta}+\frac{1}{a\sin\theta}\cdot\frac{\partial S_\theta}{\partial\varphi} \end{aligned}\right\} \tag{3}$$

Also, the tidal generating potential W_2 of the semi-diurnal tide is given by the following expression:

$$W_2 = agA\sin^2\theta\cos 2\varphi \tag{4}$$

In the case of the M_2-tide,

$$A_{M_2} = \frac{3}{2}\frac{M}{E}\left(\frac{a}{c}\right)^3\left(\frac{1}{2}-\frac{5}{4}e^2\right)\cos^4\frac{I}{2} \tag{5}$$

where M is mas of the moon, E the mass of the earth, c the average radious of the moon, e the eccentricity of the moon's orbit, I the inclination of the moon's orbit to the equotor. From Eqs. (1), (3) and (4), the components of strain are given as follows:

$$\left.\begin{aligned} e_{xx} &= 2lA\cos 2\theta\cos 2\varphi \\ e_{yy} &= -4lA\cos 2\varphi \\ e_{xy} &= -6lA\cos\theta\sin 2\varphi \end{aligned}\right\} \tag{6}$$

When the direction cosines of the direction of observations of linear strain are λ and μ, the tidal change of strain E observed will be:

$$E = e_{xx}\lambda^2 + e_{yy}\mu^2 + e_{xy}\lambda\mu \tag{7}$$

Tidal change of strain include, besides the so-called primary term, i.e., linear strain caused by tide generating force emanating from heavenly bodies, the so-called secondary term which is an indirect effect produced by the tidal changes of the sea which are caused by tide generating force of heavenly bodies.

Now, if the earth-crust in the neighborhood of the place where observations are being carried on is assumed to be an isotropic semi-infinite elastic body, strain of the earth due to the change in load of sea-water is obtained from the solution of Boussineq, as a period of the oceanic tide is long enough. When the surface of a sea-area, surrounded by concentric circles which have the radii r_n and r_{n+1}, respectively, and whose center is the spot of observations, and the movable radii making polar angles θ_n and θ_{n+1} with the direction of observations, undergoes a change of $\Delta h \cos(2t-T)$, the strain E_s in the direction of observations will be expressed as follows:

$$E_s = \frac{\rho g \Delta \cos(2t-T)}{4\pi(\lambda+\mu)} \cdot \left(\log \frac{r_{n+1}}{r_n}\right) \cdot \frac{1}{2}(\sin 2\phi_{n+1} - \sin 2\phi_n) \qquad (8)$$

provided the density of sea-water is ρ

3. Makimine is a copper mine belonging to Chichibu paleozoic system about 25km from the east coast of Kyushu. The observation room is situated in a drift about 165 m under the earth's surface. From observations of the tilting of the earth-crust in the same room Prof. Eiichi Nishimura[6] obtained very interresting results. In 1949 this place was equipped with a Sassa-type invar-line extensometer, and ever since observations have been carried on. Results of harmonic analyses of the M_2-tide observed here are shown in Table I.

Table 1 Makimine.

Location: Long. 131°24′ N.L.
Direction of observation: N 57°W.
Length of Observation Line: 20 m.
Period of Harmonic Analyses: Dec. 7, 1949—Dec. 8, 1950.
Sensibility of Instrument: $2.2\sim2.6\times10^{-9}$/mm
Observed value of M_2-tide: $(0.31\pm0.15)\times10^{-8}\cos(2t-261°\pm45°)$
Theoretical value of M_2-tide: $12.50\times10^{-8}l\cos(2t-155°)$

Complex disturbances of atmospheric origins other than the tide are seen in same specified months, so it may be best to make harmonic analyses only in months of clam weather. The value presented in this report, however was obtained from analyses made thoughout a year. Results of analyses made only in months of clam weather will be presently. The observing room at Osakayama is located 150m from earth's surface in the center of the Osakayama tunnel (700m in length) of the former Tokaido Railway Line on the outskirts of Otsu city, Shiga Prefecture. The place is equipped with extensometers three of whose constituent parts are the Sassa-type invar-line strain-meters, and the rest is a vertical component extensometer. Along with it, there is an invar-rod extensometer, for the purpose of comparative observations. The annual variation of temperature in the tunnel is 0.2°C; the daily variation is less than 10^{-2}C. The daily variation of horizontal strain is less than 10^{-9}. The neighbouring stratum is made up of clay-slate that belongs to Chichibu palaeozoic system

Table 2. Osakayama.

Location: Long. 135°51′ E. 34°59′ N.L.
I. Direction of Observations: S 38°W.
Length of Observations: 20 m.
Sensibility of Instrument: $0.41–0.63\times10$/mm.
Period of Harmonic Analyses: 18mos. from Oct. 24, 1947 to Feb. 24, 1949.
Value obtained of M_2-tide: $(0.33\pm0.10)\times10^{-8}\cos(2t-43°\pm12°)$.

Theoretical value of M_2-tide: $9.48\times10^{-8} l\cos(2t-220°)$.

II. Direction of Observations: S 76°W.
Length of Observation: 6.8 m.
Sensibility of Instrument: 2.56×10^{-8}/mm.
Period of Harmonic Analyses: from Sep. 2, to Oct. 1, 1952.
Observational Value of M_2-tide: $1.31\times10^{-8}\cos(2t-187°)$.
Theoretical Value of M_2-tide: $14.42\times10^{-8} l\cos(2t-192°)$.

III. Direction of Observations: S 76°W.
Length of Observation: 3.4 m.
Sensibility of Instrument: 5.70×10^{-9}/mm.
Period of Harmonic Analyses: from May 19, to Jun. 3, 1952.
Observational Value of M_2-tide: $0.06\times10^{-8}\cos(2t-178°)$.
Theoretical Value of M_2-tide: $2.70\times10^{-8}\cos(2t-190°)$.

Results of analyses of the M_2-tide of horizontal components in the directions of S 38°W, S 76°W and N 2°E are shown in Table 2.

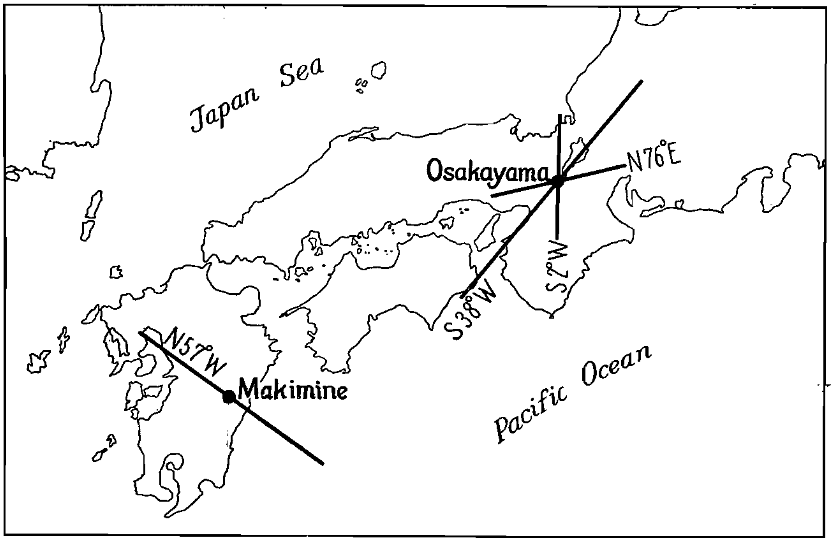
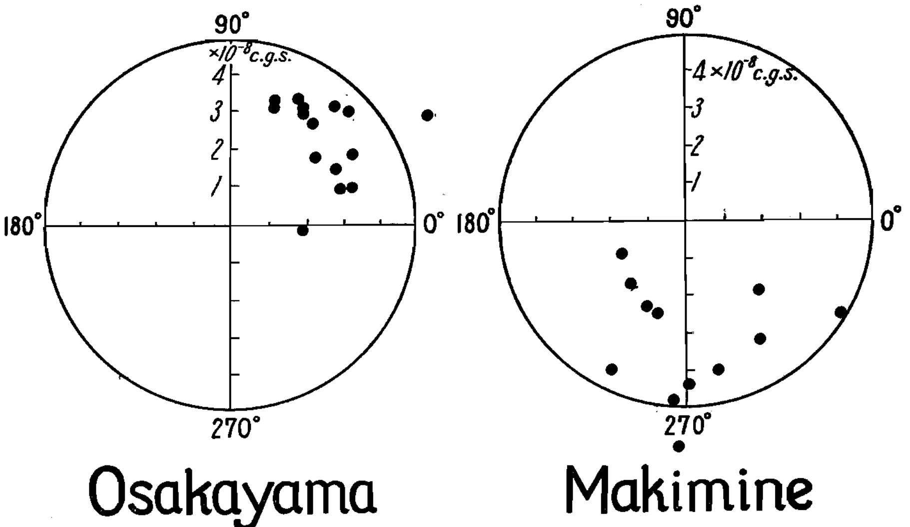


Fig. 1 Location of observatorys, and directions of observations.

In Fig. 1. are shown locations of Osakayama and Makimine, and the directions of observations. In Fig. 2. are shown observed values of the M_2-tide analysed monthly at both Osakayama and Makimine. The amplitude of strain of M_2-tide is expressed by the length of the movable radius, and the phase angle made with the initial line. As clear from those two figures, the values obtained at Osakayama concentrate in a rather small area while the values obtained at Makimine are extensively scattered around. One of the reasons for this will be that the sensibility of the instrument used at Makimine, which is 2.2–2.6 10^{-8}/mm, is only fifth to one forth of that of the instrument used at Osakayama.



Fjg. 2 Monthly obtervational value of M_2-tide.

4. What is being sought in our observations of the earth-tide by the extensometer is firstly the most probable value of the Lambert's number l, and secondly how the earth-crust is strained by the load of sea-water.

Assuming that there is no influence of the oceanic tide, it is possible to get l from the observational value and theoretical value shown above Table. 2 as following:

in the direction of S 38°W :	$l=0.035\pm0.011$	at Osakayama
in the direction of S 76°W :	$l=0.091$	at Osakayama
in the direction of N 2°E :	$l=0.021$	at Osakayama
in the direction of N 57°W :	$l=0.025\pm0.012$	at Makimine

Assuming $\lambda=\mu$ and using value $\mu=6.0\times10^{11}$c.g.s. of the depth 30–60km. from the earth's surface, horizontal strains in the direction of observations caused by the oceanic tide are calculated from Boussineq's solution is following:

$0.234\times10^{-8}\cos(2t-278°)$ at Osakayama
$0.265\times10^{-8}\cos(2t-138°)$ at Makimine

Supposing rigidity $\mu=10^{12}$c.g.s., the strain will be:

$0.139\times10^{-8}\cos(2t-278°)$ at Osakayama
$0.158\times10^{-8}\cos(2t-138°)$ at Makimine

The influence of the oceanic tide can not be ascertined independently, and consequently, it is hard to separate the observational value into the primary-term and the secondary term, unless some assumption is made, However, this question can be solved if either of two terms is given. It is not possible in a small island as Japan to observe tidal strains at a place far distant from the sea-coast. Dr. Takahiro Hagiwara[7] made observations on the contray at a place close to the seashore where the secondary term was so large that the primary term could be disregarded. As a result, it was found out that there were few instances to which the solution of Boussinesq could be applied.

Assuming in the first place various of l, values of the secondary term are obtained from the observational values, and shown in Table 3.

Table 3. Secondary term.

	Osakayama.		Makimine.	
l	Amplitude. $\times 10^{-8}$	Phase.	Amplitude. $\times 10^{-8}$	Phase.
0.04	0.71	43°	0.71	313°
0.06	0.90	43°	0.97	318°
0.08	1.09	43°	1.15	322°
0.18	2.07	43°	2.42	328°

Relations implicit in the above Table 3 are shown in Fig.

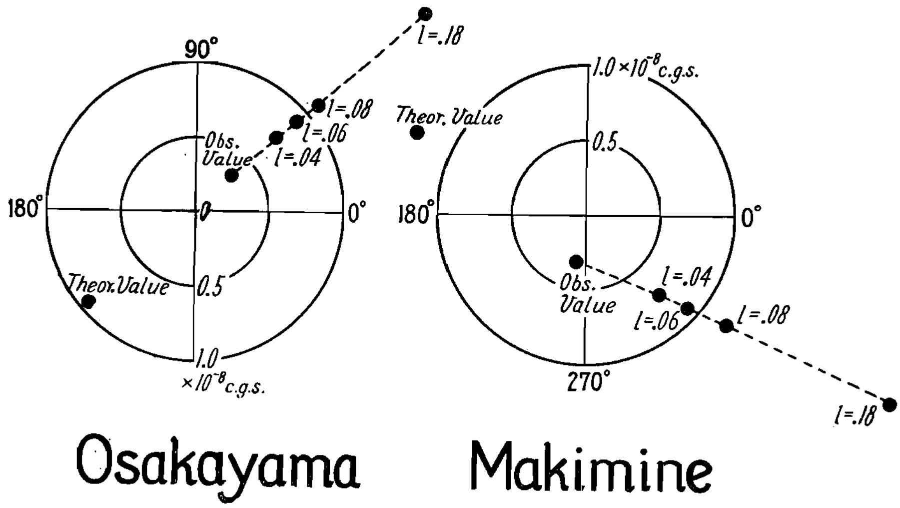


Fig. 3 Secondary-term for various values of l.

As shown in Fig. 3, in the case of Osakayama, all phase angles of the secondary term are found in the first quadrant, while in the case of Makimine, they are found in the fourth quadrant.

Eqation (8) is wirtten to the following form,

$$\left.\begin{aligned} &E_s=\psi(r_n,\ \phi_n)\cdot f(t,\ r_n,\ \phi_n)\\ &f(t,\ r_n,\ \phi_n)=\Delta h\cos(2t-T)\log\frac{r_{n+1}}{r_n}\cdot\frac{1}{2}(\sin 2\phi_{n+1}-\sin 2\phi_n)\\ &\psi(r_n,\ \phi_n)=\frac{\rho g}{4\pi(\lambda+\mu)} \end{aligned}\right\} \tag{9}$$

putting

$$f(t,\ r_n,\ \phi_n)=a_n\cos 2t+b_n\sin 2t \tag{10}$$

the value of a_n and b_n are caluculated as to a sea-ares divisible by concentric circles having the radii 25km.–40km., 40km.–63km., 63km.–100km., 100km.–160km., 160km.–250km., 250km.–400km., 400km.–630km., 630km.–880km., 880km.–1,400km., and 1,400km.–2,200km., and centering at the place of observations. They are shown in Table 4 and Fig. 4.

Table 4

	Sea-area		Osakayama		Makimine	
n	min. radius km	max. radius km	a_n	b_n	a_n	b_n
1	25	40	—	—	−5.24	—
2	40	63	−0.45	−0.25	−7.89	1.28
3	63	100	0.54	−2.76	−3.98	1.84
4	100	160	1.38	−4.38	−6.74	1.90
5	160	250	3.40	−4.89	−3.70	−0.07
6	250	400	0.34	−4.99	−1.81	−1.56
7	400	630	−0.49	−3.34	−0.04	1.55
8	630	880	−1.10	−5.84	−1.24	10.44
9	880	1400	0.66	−4.69	−1.27	4.27
10	1400	2200	−1.67	−4.36	1.30	3.98

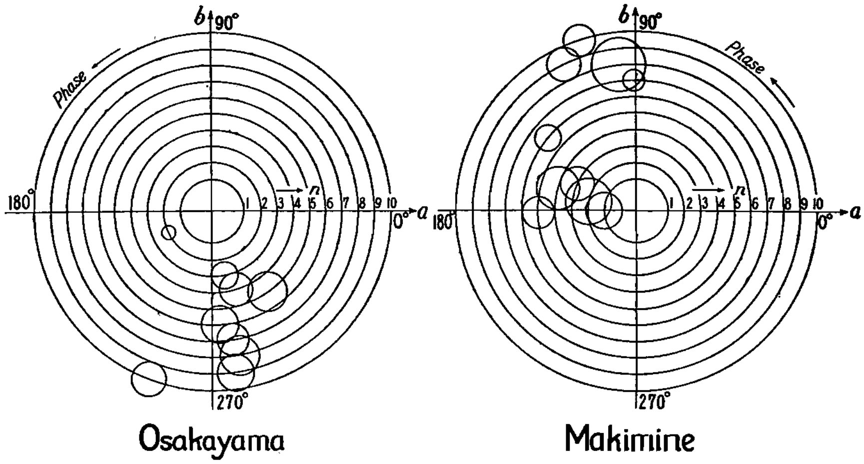


Fig. 4 $\sqrt{a^2_n+b^2_n}$ and $\theta=\tan^{-1}\frac{b_n}{a_n}$ of divided sea-area by consentric circles.

In the Figure 4, the space between the center of the large concentric circle and the center of the small circle is the distance r_n from the place of observations to the fan-shaped sea-area, the polar angle θ_n from the original line showing the position of the small circle is equal to $\tan^{-1}\frac{b_n}{a_n}$, and the area of the small circle shows the size of $\sqrt{a_n^2+b_n^2}$.

Supposing $\psi(r, \phi)$ in (9) is a function only of the distance from the observatory to the sea-area as

$$\psi(r)=\frac{1}{\alpha(1+\beta r^m)} \tag{11}$$

Next, calculations are made assuming different vslues for α and β. (In this case, a unit for large of r is 100km.) In Equation (11), the case where $\beta=0$ corresponds to the case where the earth-crust is an isotropic semi-infinite elastic body. For convenience' sake α is assumed to be 1.03×10^{10} so that $\frac{\rho g}{4\pi(\lambda+\mu)}=\frac{1}{\alpha}$ when the rigidity of the earth crust $\mu=0.4\times10^{10}$ c.g.s. and $\lambda=\mu$. Results found by calulation are shown in Table 5.

Table 5 (A)

Osakayama: $\psi=\frac{1}{\alpha(1+\beta\gamma)^m}$, $\alpha=1.03\times10^{10}$, $\beta=0.0$,

n	max. radius km	$a_i\times10^{-8}$	$b_i\times10^{-8}$	$\sum_{i=1}^{n}\psi_i a_i\times10^{-8}$	$\sum_{i=1}^{n}\psi_i b_i\times10^{-8}$	θ
1	40	——	——	——	——	——
2	63	−0.004	−0.002	−0.004	−0.002	200°
3	100	0.005	−0.027	−0.001	−0.029	289°
4	160	0.013	−0.043	0.014	−0.072	238°
5	250	0.033	−0.047	0.047	−0.119	292°
6	400	0.003	−0.048	0.051	−0.167	288°
7	630	−0.005	−0.032	0.046	−0.200	284°
8	880	−0.011	−0.057	0.035	−0.256	283°
9	1400	0.006	−0.046	0.042	−0.302	289°
10	2200	−0.016	−0.042	−0.025	−0.344	278°

Table 5 (B)

Makimide: $\psi=\frac{1}{\alpha(1+\beta\gamma)^m}$, $\alpha=1.03\times10^{10}$, $\beta=0.0$

n	max. radius km	$a_i\times10^{-8}$	$b_i\times10^{-8}$	$\sum_{i=1}^{n}\psi_i a_i\times10^{-8}$	$\sum_{i=1}^{n}\psi_i b_i\times10^{-8}$	θ
1	40	−0.051	——	−0.051	——	180°
2	63	−0.077	0.012	−0.127	0.012	175°
3	100	−0.039	0.018	−0.166	0.030	170°
4	160	−0.066	0.018	−0.231	0.048	167°
5	250	−0.036	−0.001	−0.267	0.047	170°
6	400	−0.018	0.015	−0.285	0.063	167°
7	630	−0.004	0.015	−0.285	0.078	164°
8	880	−0.001	0.102	−0.297	0.180	147°
9	1400	−0.001	0.041	−0.310	0.221	143°
10	2200	0.001	0.039	−0.297	0.260	138°

Table 5 (C)

Osakayama: $\alpha=1.03\times10^{10}$ $\beta=1.0$ $m=1$

n	$a_i\times10^{-8}$	$b_i\times10^{-8}$	$\sum_{i=1}^{n}\psi_i a_i\times10^{-8}$	$\sum_{i=1}^{n}\psi_i b_i\times10^{-8}$	θ
1	——	——	——	——	——
2	−0.003	−0.002	−0.003	−0.002	205°
3	0.003	−0.015	0.000	−0.017	270°
4	0.006	−0.019	0.006	−0.036	280°
5	0.011	−0.016	0.017	−0.047	290°
6	0.001	−0.012	0.018	−0.053	289°
7	−0.001	−0.005	0.017	−0.060	281°
8	0.001	−0.007	0.018	−0.067	280°
9	0.001	−0.004	0.019	−0.071	280°
10	−0.001	−0.002	0.018	−0.073	279°

Table 5 (D)

Makimine: $\alpha=1.03\times10^{10}$ $\beta=1.0$ $m=1$

n	$a_i\times10^{-8}$	$b_i\times10^{-8}$	$\sum_{i=1}^{n}\psi_i a_i\times10^{-8}$	$\sum_{i=1}^{n}\psi_i b_i\times10^{-8}$	θ
1	−0.039	——	−0.039	——	180°
2	−0.051	0.008	−0.090	0.008	168°
3	−0.022	0.010	−0.111	0.018	170°
4	−0.029	0.008	−0.141	0.027	169°
5	−0.012	0.000	−0.153	0.026	170°
6	−0.004	0.004	−0.157	0.030	169°
7	0.000	0.003	−0.157	0.033	168°
8	−0.002	0.012	−0.158	0.045	164°
9	−0.001	0.004	−0.159	0.048	163°
10	0.001	0.002	−0.159	0.051	162°

Table 5 (E)

Osakayama: $\alpha=1.03\times10^{10}$ $\beta=2.0$ $m=1$

n	$a\times10^{-8}$	$b\times10^{-8}$	$\sum_{i=1}^{n}\psi_i a_i\times10^{-8}$	$\sum_{i=1}^{n}\psi_i b_i\times10^{-8}$	θ
1	——	——	——	——	——
2	−0.002	−0.001	−0.002	−0.001	150°
3	0.002	−0.010	0.000	−0.012	260°
4	0.004	−0.013	0.004	−0.024	285°
5	0.007	−0.010	0.010	−0.034	288°
6	0.001	−0.007	0.011	−0.041	286°
7	0.000	0.003	0.010	−0.043	284°
8	0.001	−0.004	0.011	−0.047	284°
9	0.000	−0.002	0.011	−0.049	284°
10	−0.001	−0.001	0.011	−0.050	283°

Table 5 (F)

Makimine: $\alpha = 1.03 \times 10^{10}$ $\beta = 2.0$ $m = 1$.

n	a_i	b_i	$\sum_{i=1}^{n} \psi_i a_i \times 10^{-8}$	$\sum_{i=1}^{n} \psi_i b_i \times 10^{-8}$	θ
1	−0.031	——	−0.031	——	180°
2	−0.038	0.006	−0.069	0.006	175°
3	−0.015	0.007	−0.084	0.013	166°
4	−0.019	0.005	−0.103	0.019	160°
5	−0.007	0.000	−0.110	0.019	161°
6	0.002	0.002	−0.113	0.021	160°
7	0.000	0.001	−0.114	0.029	160°
8	−0.001	0.007	−0.114	0.030	156°
9	−0.001	0.002	−0.114	0.030	155°
10	0.000	0.001	−0.114	0.032	155°

Table 5 (G)

Osakayama: $\alpha = 1.03 \times 10^{10}$ $\beta = 0.5$ $m = 2$

n	$a_i \times 10^{-8}$	$b_i \times 10^{-8}$	$\sum_{i=1}^{n} \psi_i a_i \times 10^{-8}$	$\sum_{i=1}^{n} \psi_i b_i \times 10^{-8}$	θ
1	0.000	0.000	0.000	0.000	——
2	−0.004	−0.002	−0.004	−0.002	206°
3	0.004	−0.020	0.000	−0.022	270°
4	0.008	−0.024	0.008	−0.046	280°
5	0.011	−0.016	0.019	−0.062	286°
6	0.001	−0.008	0.020	−0.070	286°
7	−0.000	−0.002	0.019	−0.072	285°
8	0.000	−0.002	0.020	−0.075	285°
9	0.000	−0.001	0.020	−0.075	285°
10	0.000	0.000	0.020	−0.075	285°

Table 5 (H)

Makimine: $\alpha = 1.03 \times 10^{10}$ $\beta = 0.5$ $m = 2$

n	$a_i \times 10^{-8}$	$b_i \times 10^{-8}$	$\sum_{i=1}^{n} \psi_i a_i \times 10^{-8}$	$\sum_{i=1}^{n} \psi_i b_i \times 10^{-8}$	θ
1	−0.050	——	−0.050	——	180°
2	−0.066	0.011	−0.114	0.011	175°
3	−0.029	0.014	−0.143	0.024	170°
4	−0.036	0.010	−0.180	0.035	169°
5	−0.012	−0.000	−0.192	0.035	170°
6	−0.003	0.003	−0.195	0.037	170°
7	0.000	0.001	−0.195	0.038	169°
8	−0.000	0.004	−0.195	0.042	168°
9	−0.000	0.001	−0.196	0.043	168°
10	0.000	0.000	−0.196	0.043	168°

Table 5 (I)
Osakayama : $\alpha=1.03\times10^{10}$, $\beta=1.0$ $m=2$

n	$a_i\times10^{-8}$	$b_i\times10^{-8}$	$\sum_{i=1}^{n}\psi_i a_i\times10^{-8}$	$\sum_{i=1}^{n}\psi_i b_i\times10^{-8}$	θ
1	——	——	——	——	——
2	−0.003	−0.002	−0.003	−0.002	213°
3	0.003	−0.016	0.000	−0.018	270°
4	0.005	−0.016	0.005	−0.035	278°
5	0.007	−0.010	0.012	−0.044	286°
6	0.000	−0.004	0.012	−0.048	285°
7	0.000	−0.001	0.012	−0.050	284°
8	0.000	−0.001	0.012	−0.051	284°
9	0.000	−0.000	0.012	−0.051	284°
10	−0.000	−0.000	0.012	−0.051	284°

Table 4 (J)
Makimine : $\alpha=1.03\times10^{10}$, $\beta=1.0$ $m=2$

n	$a_i\times10^{-8}$	$b_i\times10^{-8}$	$\sum_{i=1}^{n}\psi_i a_i\times10^{-8}$	$\sum_{i=1}^{n}\psi_i b_i\times10^{-8}$	θ
1	−0.046	——	−0.046	——	180°
2	−0.058	0.010	−0.105	0.010	174°
3	−0.024	0.011	−0.128	0.020	171°
4	−0.025	0.007	−0.153	0.028	170°
5	−0.007	0.000	−0.161	0.028	170°
6	−0.002	0.001	−0.162	0.029	170°
7	0.000	0.002	−0.163	0.031	169°
8	0.000	0.000	−0.163	0.031	169°
9	0.000	0.000	−0.163	0.031	169°
10	0.000	0.000	−0.163	0.031	169°

Table 5 (K)
Osakayama : $\alpha=1.03\times10^{10}$, $\beta=2.0$ $m=2$

n	$a_i\times10^{-8}$	$b_i\times10^{-8}$	$\sum_{i=1}^{n}\psi_i a_i\times10^{-8}$	$\sum_{i=1}^{n}\psi_i b_i\times10^{-8}$	θ
1	——	——	——	——	——
2	−0.003	−0.002	−0.003	−0.002	213°
3	0.002	−0.013	−0.001	−0.015	270°
4	0.003	−0.013	0.002	−0.028	275°
5	0.004	−0.004	0.006	−0.032	281°
6	0.000	0.000	0.006	−0.032	281°
7	0.000	−0.001	0.006	−0.033	281°
8	0.000	−0.000	0.006	−0.033	281°
9	0.000	0.000	0.006	−0.033	281°
10	0.000	0.000	0.006	−0.033	281°

Table 4 (L)

Makimine : $\alpha=1.03\times10^{10}$, $\beta=2.0$ $m=2$

n	$a_i\times10^{-8}$	$b_i\times10^{-8}$	$\sum_{i=1}^{n}\psi_i a_i\times10^{-8}$	$\sum_{i=1}^{n}\psi_i b_i\times10^{-8}$	θ
1	−0.043	——	−0.043	——	180°
2	−0.059	0.010	−0.102	0.010	175°
3	−0.017	0.008	−0.119	0.018	171°
4	−0.016	0.004	−0.135	0.022	170°
5	−0.004	0.000	−0.139	0.022	169°
6	−0.001	0.001	−0.140	0.023	170°
7	0.000	0.000	−0.140	0.023	170°
8	0.000	0.001	−0.140	0.024	170°
9	0.000	0.000	−0.140	0.024	170°
10	0.000	0.000	−0.140	0.024	170°

Fig. 5 (a) $\psi(\gamma)\cdot f(\gamma,\phi)$ at Osakayama.

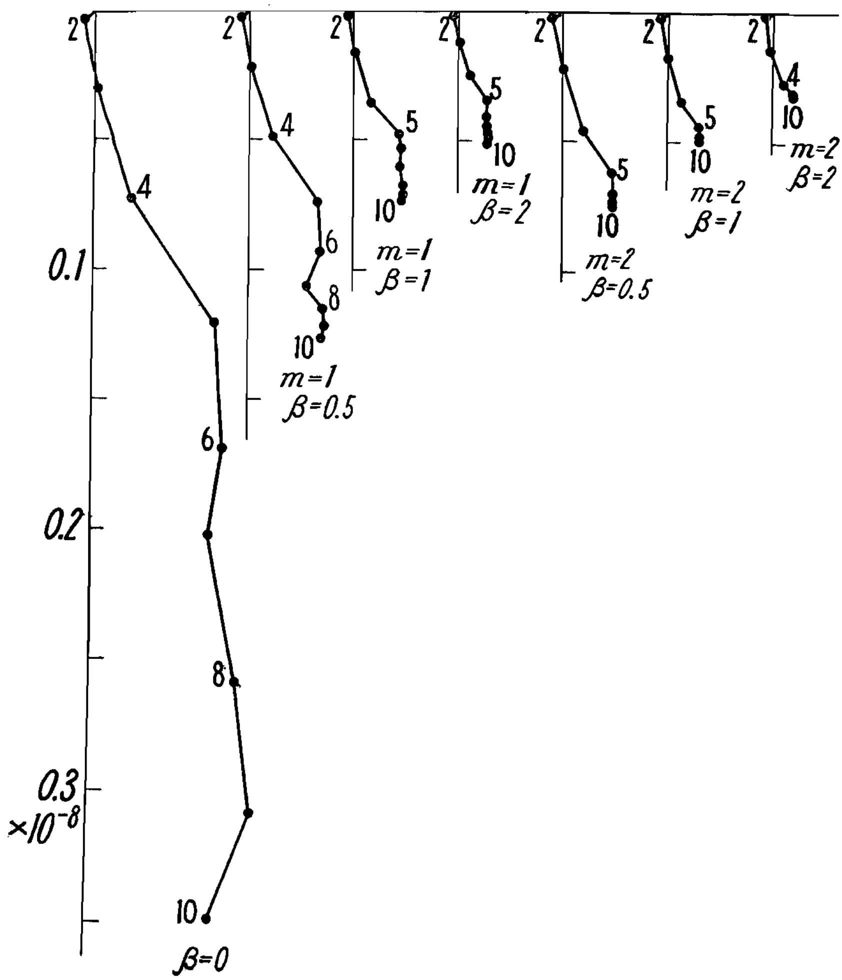
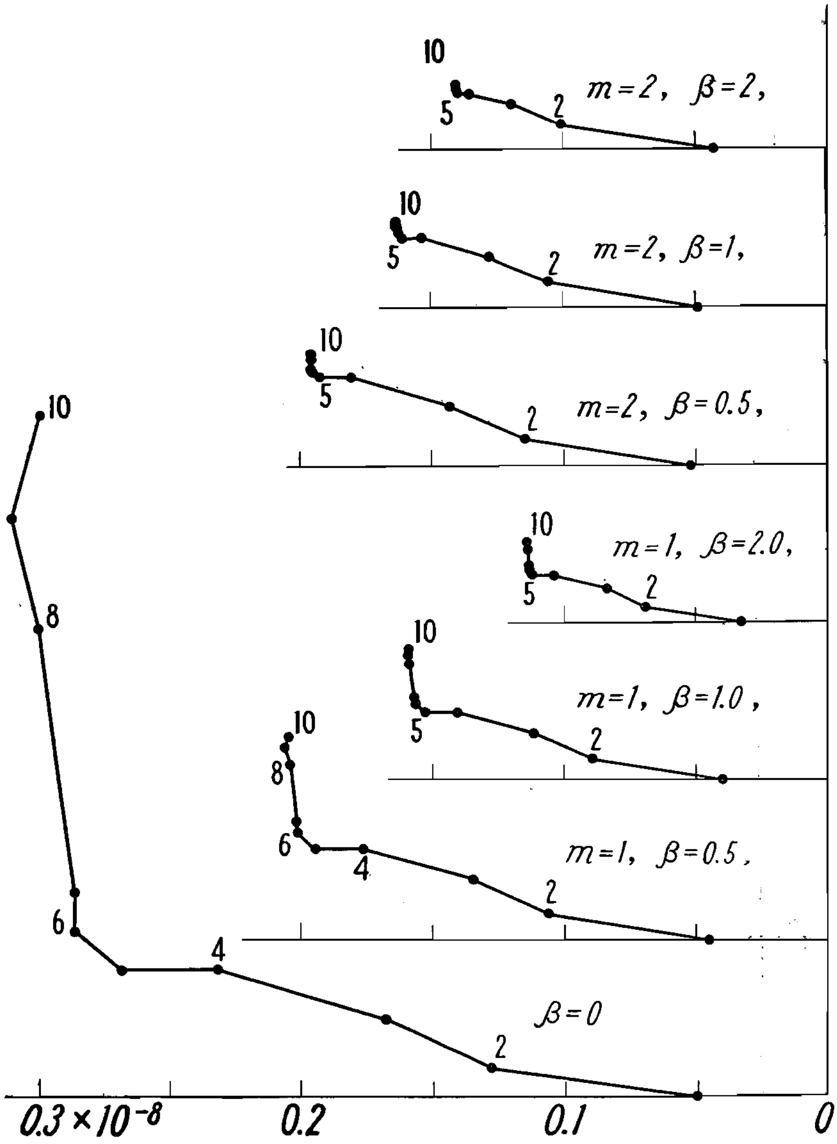


Fig. 5 (b) $\psi(\gamma)\cdot f(\gamma,\phi)$ at Makimine.



Besides the above Table 5, in Fig. 5 some case $m=1$, 2, and $\beta=0.5$, 1.0, 2.0, are graphically illustrated. As clear from the above Table and Fig. 5 and also from Fig. 4, it never happens that the value of $\theta_n=\tan^{-1}\frac{\sum_{i=1}^{n} b_i}{\sum_{i=1}^{n} d_i}$ is found in the first quadrant in the case of Osakayama, and in the fourth quadrant in the case of Makimine, so long as assume $\psi(r)$ as in a form of (11) moreover amplitude is very much diflerent. Therefore, it is not proper for us to suppose that $\psi(r,\phi)$ is a form of (11). It is very interesting that however, is that according to Table, when $\psi(r,\phi)=$const., the observational value is out of phsse to calculated value in the secondary term by 130° at Oakayama, and 180° the case of Makimine. This means probablly that our assumption in this calculation is not satisfied in full. It is necessary to make observation at least in three directions at one place even of horizantal strain alone. Not withstanding in foregoing pages we discussed the results of observations made at Osakayama and Makimine in only one direction. Observations of three hrizontal components are carried on at Osakayama. Therefore more extended discussion will be reported in the next occation waitting accumlation of data of strain observations of three horizontal components which are now carried on Osakayama and others. Our cordinal thanks are extended to Dr. Kenzo Sassa, Professor of Kyoto University who instructed us in these studies, and also to Prof. Eiichi Nishimura who furnished us with the records of observations at Makimine.

Disaster Prevention Research Institute
Kyoto University.
1952 August 1.

References

(1) Sassa, K., Ozawa, I. and Yoshikawa, S.; Observation of Tidal Strain of the Earth. International Association of Geodesy. Brusells Assembly, August 1951.
Sassa, K. Ozawa, I. and Yoshikawa, S,; Observstion of Earth-tide by New Horizontal Extensometer. Bulletin of the Disaster Prevention Research Institute Kyoto University No. 2. (1949). pp. 113.

(2) Nishimura, E.; On the Eaath-tide part III. " Tikyubutsuri." vol. 5. 1941. pp. 89.
Nishimura, E.; On the Earth-tide Transactions American Geophysical Union vol. 31. Part I. pp. 357.

(3) Nishimura, E.; On the Earth-tide Part IV. "Tikyubutsuri" vol. 5. 1941 pp. 113.

(4) Pollak, L.W.; Das Periodgramme per Polbewegung. Beiträge zur Geophsik. Band 16. 1927.
Lamor, J.; On the Influence of the Oceanic Waves on the Law of Variation of Latitude. Proceeding of London Mathematical Society vol. 14. pp. 440.

(5) Takeuchi, H.; On the Earth-tide "Kagaku" vol. 18. 1948. pp. 60.
Takeuchi, H.; Theoretical study of Earth-tide. Transactions Geophysical Union vol. 31. (1950) No. 5. Part I. pp. 651.

(6) Nishimura, E.; On the Earth-tide Part VI. (1947) Report of Geophysical Institute Kyoto Imperial University.

(7) Hagiwara, T., Rikitake, T., Kasahara, K., & Yamada, J.; Observation of the Deformation of the Earth's Surface at Aburatsubo, Miura Peninsura. Part III.. Bulletin of Earthquake Research Institute Vol. 27. (1949). pp. 39.

(8) Ogura, S.; On the Tide of the Japanese neighbouring Sea. Bulletin of Hydrographic Depertment Imperial Japanese Navy vol. VII. (1938).